

TABLE 6. Cohesite Parameters for Various Cases

Case	$K_0$	$M_b$	$K_0^*$	$(ak_0/3\pi)^{2/3}$	$y$	$d \ln y$	$d \ln V$	$d \ln \bar{V}$
1	1.27	5.6	-0.05	0.43	-0.04	4.9		
2	1.36	4.1	-0.05	0.46	1.2	4.6		
3	0.974	7.3	-0.05	0.33	-0.15	6.4		

SiO<sub>2</sub> PHASE EQUILIBRIA

They cannot definitely be said to be disordrant with case 3, but the discrepancy between case 3 and the static-compression data seems to be significant. Because of this discrepancy, the equation of state of coesite must remain some-what uncertain at this stage.

Data Error, Mb	S11 0.20	S12 0.10	S13 0.10	X3 0.02
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TABLE 5. Standard Errors Assumed

Initially both  $K_a$  and  $K_c$ , were determined by the Hugo mot and static-compression data. The results are given as cases 1 and 2 in Table 6, case 1 excluding the three doubtful Hugo mot points and case 2 including them. The standard errors used to weight the compression data are given in Table 5. Case 1 is illustrated in Figure 1, case 2 in Figure 6. The bulk moduli in these two cases are significantly above the value of 0.97 N/m measured ultrasonically by Mizutani et al. (H. Mizutani, private communication, 1972), and so a third case was run with  $K_a = 1972^2$ , and at this value and only  $K_c$ , determined by the compression data (Table 6 and Figure 6). It can be seen (Figure 6) that case 3 does not fit the static-compression data of Bassett and Barnesett [1970] very well, and it fails below most of the corresponding Hugo mot data.

It can be seen from Figure 1 that the  $p_0 = 1.35 \text{ g/cm}^3$  Hugoniot data are considerably scattered and that they do not trend toward the core site density of  $2.91 \text{ g/cm}^3$ , perhaps because there has been a partial conversion to the stoichiometric phase. When they are compared to the  $p_0 = 1.15 \text{ g/cm}^3$  Hugoniot data, the lower three points in particular are seen to deviate toward higher densities. Two cases were therefore treated, one including these three points, the other excluding them.

not strongly affect the equation of state in this range. A value of  $-0.05 \text{ kJ/K}$  was therefore assumed. This value of  $(\partial K/\partial T)^p$  gives values of  $\delta_p$  in the range 5–10, a range that seems reasonable on the basis of a few other examples, including stishovite [e.g., Anderson *et al.*, 1998; Roberts and Ruppel, 1971]. The values of  $V$ , and  $a$  were taken from Table 2, and  $C_p$  was calculated from the Debye model.

Because of the smaller range and quantity of 'coarse' data, it is not possible to determine as many parameters of the equation of state as were determined for the isobaric data. Because the data extended to only about 15% volume compression, it is not necessary to use the full fourth-order version of (3), and so the energy term is here assumed to be zero. Because there is not a large range in the initial porosities of the Hugoiont data, the volume dependence of  $(\partial k / \partial T)^p$  cannot be well determined from the data.

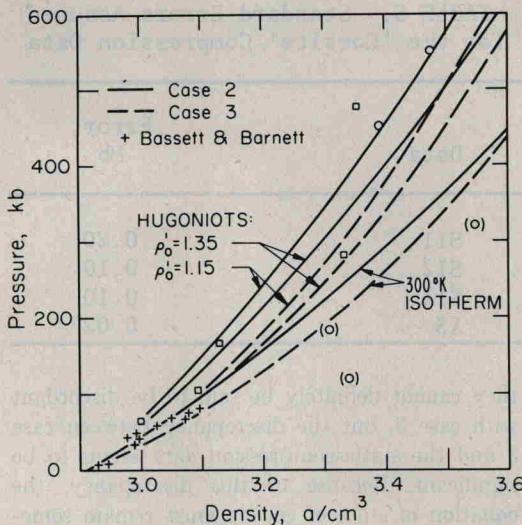


Fig. 6. Hugoniot data of 'coesite' and calculated Hugoniots and 300°K isotherms from cases 2 and 3 (Table 6). Symbols are those used in Figures 1 and 5.

the points to move along the Hugoniot locus, which in a  $P$ - $T$  plot is approximately radial from the initial point.

The boundary between the 'coesite' and stishovite fields (Figure 8) is closely defined by the  $\rho_0' = 1.77$  and  $\rho_0' = 1.55 \text{ g/cm}^3$  Hugoniot points, both of which show signs involving a mixture of the two phases, as was discussed earlier.

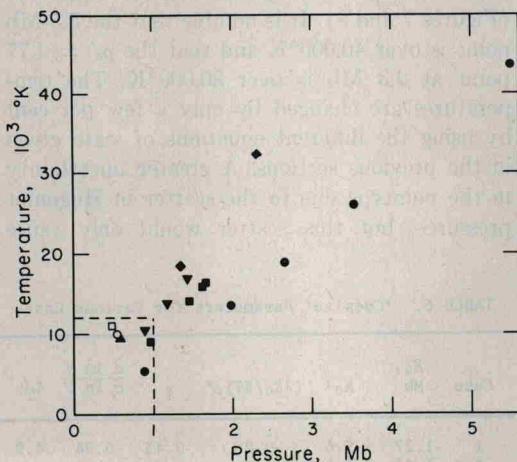


Fig. 7. Calculated Hugoniot temperatures of stishovite and 'coesite' versus Hugoniot pressure. Box is enlarged in Figure 8. Symbols are those used in Figure 1.

The Gibbs free energy is defined by

$$G = H - TS = U + PV - TS \quad (14)$$

where  $H$  is the enthalpy and  $S$  is the entropy. Here  $G$  has the property [e.g., Slater, 1939]

$$(\partial G / \partial P)_T = V \quad (15)$$

We wish to evaluate  $G$  at the state  $(P, V, T)$ , starting from the state  $(V_0, P_0, T_0)$ . (Atmospheric pressure can be ignored here.) This evaluation will be done via the state  $(P_0, V_0, T)$ , where  $P_0(T) = P(V_0, T)$  (i.e., by first raising the temperature at constant volume and then compressing isothermally). From (14)

$$\begin{aligned} G(V_0, T) &= G(V_0, T_0) \\ &+ [U(V_0, T) - U(V_0, T_0)] + P_0(T) V_0 \\ &- [TS(V_0, T) - T_0 S(V_0, T_0)] \end{aligned} \quad (16)$$

and from (15), upon integration,

$$G(V, T) = G(V_0, T) + \int_{P_0(T)}^{P(T)} V(P', T) dP' \quad (17)$$

When the difference between the Gibbs free energies of stishovite and coesite at the state  $(V_0, T_0)$  are denoted by  $\Delta G_0$  (i.e.,

$$\Delta G_0 = G^s(V_0, T_0) - G^c(V_0, T_0)$$

where superscripts  $s$  and  $c$  denote stishovite and

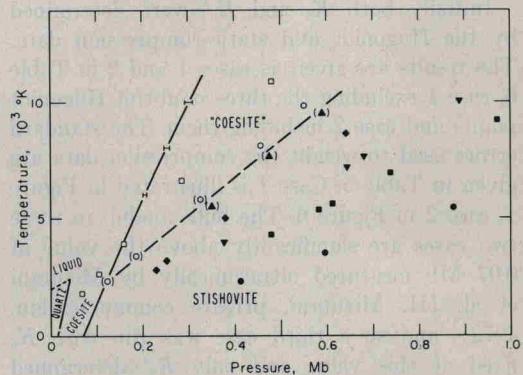


Fig. 8. Calculated Hugoniot temperatures of stishovite and 'coesite' versus Hugoniot pressure compared with observed and calculated (solid and short-dashed) phase lines. Long-dashed line separates stishovite and 'coesite' fields. Error bars represent variations due to the use of alternative equations of state given in previous sections. Symbols are those used in Figure 1.